

Elastic-acoustic coupling for large scale earthquake simulations

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Earthquake-Tsunami Coupling Workflows

One-way Linked



Figure from: H. Madden et al. "Linked 3-D modelling of megathrust earthquake-tsunami events: from subduction to tsunami run up (2021)

Fully Coupled









One-way linking vs 3D coupling

Using shallow water equations for tsunami has disadvantages:

- No dispersion (if not using Boussinesq approximation)
- No acoustic waves (i.e., assuming incompressible ocean) -> Potentially dominant in data recorded by offshore instruments
- Only works in **shallow water** limit

Fully-coupled elastic-acoustic model solves entirely new class of earthquake-tsunami problem

Compares well with one-way linking under certain conditions

Detailed model comparison: Abrahams, Lauren S., et al. "Comparison of methods for coupled earthquake and tsunami modelling." *Geophysical Journal International* 234.1 (2023)



Example: Palu, Sulawesi September 2018

- Mw 7.5 strike-slip earthquake Propagation at supershear speed crossing narrow Palu Bay
- Followed by unexpected and localized tsunami
- Complicated geometry: bath-tub like bay, very shallow water (average 600 m)
- Details: L. Krenz et al. "3D Acoustic-Elastic Coupling with Gravity: The Dynamics of the 2018 Palu, Sulawesi Earthquake and Tsunami". Proceedings of the International Conference for High Performance Computing, Networking, Storage and Analysis, 2021





SeisSol

What

- (An)Isotropic elastic wave propagation
- Acoustic wave propagation
- Viscoelastic wave propagation
- Poroelasticity
- Off-fault plasticity
- Dynamic earthquake rupture How
- Numerics: ADER-DG
- Unstructured tetrahedral meshes with local time-stepping
- Optimized Hybrid MPI + OpenMP Parallelization

Available (open-source) at https://github.com/SeisSol/SeisSol/





Ocean Model



Linear acoustic medium, q = (u, v, w, p) Treated as special case of elastic wave equation with stress tensor $\sigma_{ij} = -p \ \delta_{ij}$, density ρ and $g = 9.81 \frac{m}{s^2}$.

Free surface

 $p(x, y, \eta) = 0$

Typically solved by moving mesh. Following (Lotto, Dunham 2015),linearized to: $p(x, y, z = 0) = \rho g \eta(x, y)$ $\frac{\partial \eta}{\partial t} = v_z$



Palu: Our setup

- Added water layer to existing earthquake model (Ulrich et al., 2019).
- Fully coupled model (including plasticity, seismic and acoustic waves, dynamic, earthquake rupture)
- Two meshes: M (89 million elements), L (518 million elements)
- Poly. Order 5, 46 and 261 billion degrees of freedom
- M: 5.3 hours on 1000 nodes of SuperMUC-NG for 100s simulated time
- L: 5.5 hours on 3072 nodes of SuperMUC-NG for 30s simulated time





Palu: 3D View at 15s





Comparison with One-Way Linking

Left: One-way linking, right: fully-coupled

Shows: Sea surface displacement for t = 1s, ..., 100s

Use identical earthquake model

Matches well overall, with some differences (e.g. at coast, "smoother" tsunami)





Strong Scaling on SuperMUC-NG



"Big blast followed by a long 10-second echo." Helsinki, 2018-07-08 20:37

Enhanced Geothermal System in Helsinki

- Otaniemi project
- Enhanced geothermal system (EGS), stimulated in June and July 2018 in the region of Helsinki
- Thousands of **induced**, **small earthquakes**
- Observations of ground shaking and audible disturbances collected by Macroseismic questionnaire of the Institute of Seismology, University of Helsinki
- More details: Krenz, Lukas, et al. "Numerical simulations of seismo-acoustic nuisance patterns from an induced M1. 8 earthquake in the Helsinki, southern Finland, metropolitan area." *arXiv preprint arXiv:2211.03647* (2022).





Induced Earthquakes & Reports







PGV and SPL maps SPL Horizontal SPL SPL P Wave PGV S Wave Epicentral Northings [km] 0 0 0 0 -4 Δ **Epicentral Epicentral** Epicentral Epicentral Eastings [km] Eastings [km] Eastings [km] Eastings [km] 0.0 0.05 0.1 0.02 0.045 0.0 0.01 0.02 0.0 0.02 0.045 0.0 peak horizontal peak SPL [Pa] P-wave S-wave peak SPL [Pa] peak SPL [Pa] velocity [mm/s]

PGV: Horizontal Peak Ground Velocity

SPL: Sound pressure level, reconstructed from peak vertical velocity

Conclusion

- Fully coupled elastic-acoustic simulations capture more effects than typical one-way linking strategies
- Linearization of free surface boundary conditions efficient way of tracking sea surface height
- Pronounced differences in Palu scenario: "smoother" tsunami
- Differences will be important when connecting to tsunami observations
- Further application: Modeling **sound** generated by **induced earthquakes** (due to geothermal energy)
- Outlook: Fully-coupled models for Mediterranean tsunami (Hellenic arc), Húsavík-Flatey Fault Zone, North Iceland